

# Monocacy Creek Streamflow Modeling Project

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## INTRODUCTION

Climate change is expected to result in more precipitation, and potentially more flooding, in the Lohigh Valley. In this project, I developed a hydrological model and set up of the Monocacy Creek to quantify potential future changes in flooding. Using HEC-HMS, a widely-used, publicly available hydrologic model, I first divided the Monocacy Creek into three sub-basins, allowing for some spatial variability in parameters. I selected 18 flooding events in the basin and associated the storm rainfall to have a catalog of storm events with hourly streamflow observations upstream where I collected and validated the model. In a series of sensitivity experiments, I found the model to be most sensitive to the curve number (which account for surface effects on flow), impervious surface, and canopy cover. Increasing the curve number or the impervious surface results in more streamflow, while increasing canopy cover reduces the streamflow. Ongoing work is focused on calibrating the model to match historical flood peaks, volumes, and duration. After this is completed, I will use precipitation data from future climate model scenarios to determine how flooding may change in the future. The Monocacy Creek project will be challenging to model, as a portion lies over large areas of forests, which results in significant streamflow loss to groundwater.



## HEC-HMS MODEL EXAMPLE (FIGURE 1)

Basin cells and river stream segments are represented by polygons and lines, respectively. The HEC-HMS model segments are shown as polygons and lines, respectively. The HEC-HMS model segments are shown as polygons and lines, respectively. The HEC-HMS model segments are shown as polygons and lines, respectively.

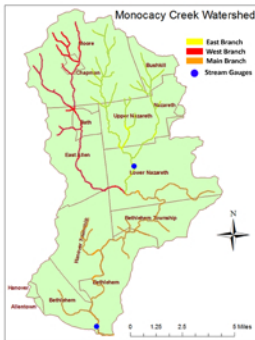
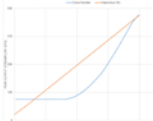


FIGURE 2 - MONOCACY CREEK WATERSHED MAP  
Image sourced from the Watershed Outlines of the Lohigh Valley by permission.

## PARAMETER SENSITIVITY TESTING FOR CALIBRATION

FIGURE 3 - CURVE NUMBER PARAMETERS



$$P_t = \frac{(P - I_a)^2}{P - I_a + S}$$

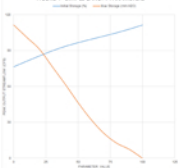
- $P_t$  = Accumulated precipitation excess at time  $t$
- $P$  = Accumulated rainfall depth at time  $t$
- $I_a$  = Initial loss
- $S$  = Potential maximum retention of soil

## CANOPY STORAGE

Canopy interception represents precipitation that is captured on trees, shrubs, and grasses, and does not reach the soil surface. Once after the storage is filled, direct precipitation becomes available for filling other storage volumes such as soil or the waterway itself.

The parameters are in two separate units. I varied both between 0 and 100. The initial storage being 0 to 100 percent and the max storage being 0 to 100 percent of the initial storage.

FIGURE 4 - SIMPLE CANOPY PARAMETERS



## MANNING'S N EQUATION FOR CHANNEL ROUTING

$$V = \frac{CR^{2/3}S^{1/2}}{n}$$

- $V$  = Average velocity
- $R$  = Hydraulic radius (ratio of channel cross section area to wetted perimeter)
- $S$  = Slope of energy grade line
- $C$  = Conversion constant (2.48 in SI case)
- $n$  = Manning's variable

FIGURE 5 - APRIL 2011 STORM RUN OUTPUT



The model run shows in this graph is early in the calibration process as the parameters are added in, the model prediction results that do not align well with the actual data. As the model development progresses, the graph alignment will improve. The model uses a Manning's  $n$  routing method for an in-channel routing and a curve number for runoff routing. The storm used for this run is highlighted in the precipitation graph table.

## THE DATA ISSUE:

The USGS (United States Geological Survey) provides access to data from thousands of stream gauges throughout the country. However, about 0.37% of the time the gauge records a data point twice. I developed a program to remove the repeated data and average out the 15 minute data to hourly, which is the required resolution of HEC-HMS.

To determine which streamflow values caused flooding of the Monocacy, I matched historical flood records to storm events as identified in historical Lohigh Valley newspapers, and then searched for the relevant streamflow records.

Table 1: Storm Data from Allentown Rain Gauge

Sourced from <a href="http://weather.gov">weather.gov</a>	
Date	Total Rainfall (mm)
01/19/96	87.818
05/17/02	35.48
07/01/07	64.523
12/10/08	64.262
04/01/09	66.428
06/17/09	62.818
08/02/09	36.83
10/24/09	58.354
12/02/09	49.784
04/03/11	66.99

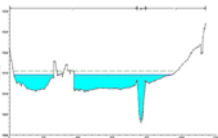
## STORM DATA:

These storms were chosen due to them being relatively recent reflecting the modern urbanized watershed. Storms are of relatively similar magnitude facilitating storm comparisons. In the future, some of these storms will be used for calibration and others for validation of the calibrated model. Moving forward some of the older storms may be required with more recent events as more data are collected.

## Future Work

### HEC-RAS CROSS SECTION (FIGURE 6)

A simple HEC-RAS cross section model with streamflow. HEC-RAS is an one-dimensional model that calculates water flow through streams and ditches and other channels.



## CALIBRATION AND VALIDATION

Calibration uses observed hydrogeomorphic data in a systematic search for parameters that yield the best fit of the computed results to the observed runoff. This search is often referred to as "optimization."

The next step is to break the entire watershed down into subbasins and build model sections of the creek using HEC-HMS.

- 1) Calibrate HEC-HMS for each subbasin until I have done the sensitivity
- 2) Validate HEC-HMS using different storms than the ones I use for validation
- 3) Create cross sections and routing HEC-HMS
- 4) Apply future precipitation data from climate models to the finished model